

Latest Paleocene deep-sea extinction: A pre-anthropogenic super greenhouse?

Ellen Thomas, Department of Earth and Environmental Sciences, Wesleyan University

Deep-sea benthic foraminifera (unicellular organisms) exhibit slow evolutionary turnover, and did not suffer during the mass extinction at the end of the Cretaceous. But about 55.5 million years ago 30-50% of their species became extinct over a few thousand years [Thomas & Shackleton, 1995]. Worldwide, post-extinction faunas show evidence of carbonate dissolution (as do bottom-dwelling crustaceans — ostracodes; Steineck & Thomas, [1996]), as well as of low oxygen conditions in many regions [Thomas, in press]. During the extinction global deep waters and high-latitude surface waters warmed by 3-4°C, remaining warm for 50-200 thousand years. Carbon isotope values of dissolved carbonate in surface and deep ocean waters and CO₂ in the atmosphere decreased by $\pm 2\text{‰}$ [Kennett & Stott, 1991; Bralower et al., 1995]. Such a large carbon isotope anomaly can not have been caused by destruction of the land biomass (no evidence in fossils) or by injection of volcanogenic CO₂; it was so rapid that it can not have been caused by erosion of organic-rich sediment. A speculative scenario: volcanogenic CO₂ (North Atlantic Volcanic Province) warmed high latitudes, where surface waters reached a low density and could no longer sink to form intermediate-deep waters. They were replaced by warmer, salty waters derived from subtropical latitudes: the deep sea warmed, leading to dissociation of (carbon isotopically very light) methane hydrates [Dickens et al., 1995]. Methane was oxidized, causing

lower oxygen, carbonate corrosive conditions in the oceans and additional warming. Benthic foraminifera became extinct as a result of low oxygen, carbonate-corrosive conditions, as well as by possibly changes in locations of high productivity (upwelling) resulting from the changes in deep-water circulation [Thomas, in press]. ODP sites thus have given us information that short-term events suggestive of major climate instabilities can occur during warm periods of Earth history; additional research (e.g., Legs 165, 167) is expected to increase our insight in this extraordinary period.

References:

- Bralower, T. J., J.C. Zachos, E. Thomas, M.N. Parrow, C.K. Paull, D.C. Kelly, I. Premoli-Silva, and W.V. Sliter, Late Paleocene to Eocene Paleoceanography of the Equatorial Pacific Ocean: Stable Isotopes recorded at ODP Site 865, Allison Guyot, *Paleoceanography*, **10**, 841-865, 1995.
- Dickens, G. R., J.R. O'Neil, D.K. Rea, and R.M. Owen, Dissociation of oceanic methane hydrate as a cause of the carbon excursion at the end of the Paleocene, *Paleoceanography*, **10**, 841-865, 1995.
- Kennett, J. P., and L.D. Stott, Abrupt deep-sea warming, paleoceanographic changes, and benthic extinctions at the end of the Palaeocene, *Nature*, **353**, 319-322, 1991.
- Thomas, E., and N.J. Shackleton, The Palaeocene-Eocene benthic foraminiferal extinction and stable isotope anomalies, *Geological Society London, Special Publication*, **101**, 401-441, 1996.
- Steineck, P. L., and E. Thomas, The latest Paleocene crisis in the deep-sea: ostracode succession at Maud Rise, Southern Ocean, *Geology*, **24**, 583-586, 1996.
- Thomas, E., The biogeography of the late Paleocene benthic foraminiferal extinction, In: M.-P. Aubry, S. Lucas, and W. A. Berggren, eds., Late Paleocene-early Eocene Biotic and Climatic Events in the Marine and Terrestrial Records, *Columbia University Press*, in press.

